

# Multi-event analysis of Lidar and Sunphotometer observations acquired at a site in Southern Canada (steps towards a regional climatology)

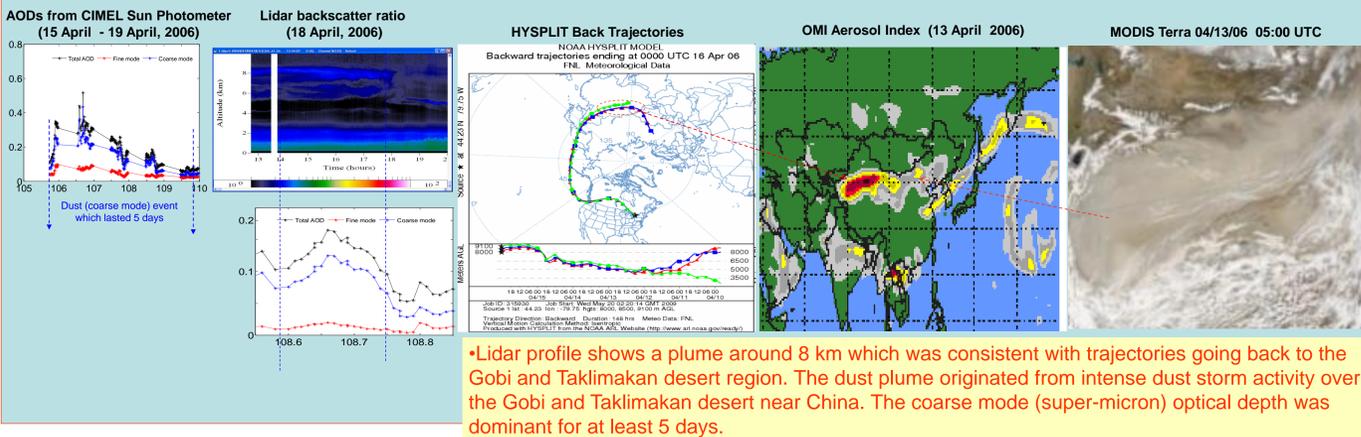
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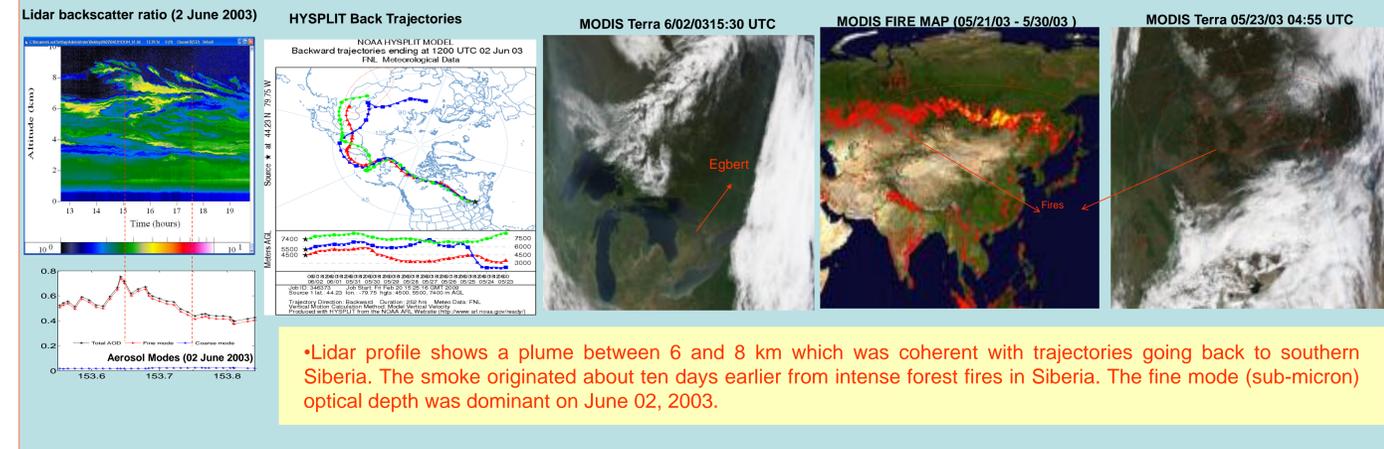
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**Objective:** To characterize regional aerosols at a rural / urban site (Egbert, Ontario) using active and passive remote sensing techniques. This is part of a more long term objective to characterize the key link between sunphotometry AODs and lidar backscatter coefficient profiles, the lidar (backscatter to extinction) ratio ( $S_a$ )  
**Instruments used:** CIMEL Sunphotometer (AEROCAN/AERONET network), ALIAS (Aerosol Lidar for Atmospheric Studies)  
**Remote sensing imagery products used:** MODIS (Aqua and Terra), Aerosol Index (OMI), HYSPLIT Back Trajectories  
**Models employed:** HYSPLIT (Hybrid Single Particle Lagrangian Integrated Trajectory Model)

## 1. Long range transport of dust aerosols



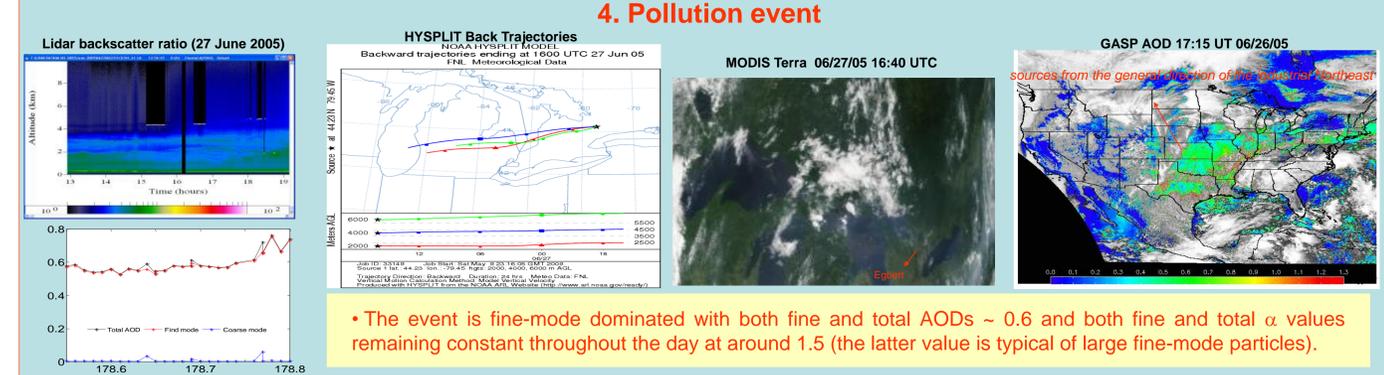
## 2. Long range transport of smoke aerosols



## 3. Regional transport of smoke aerosols



## 4. Pollution event



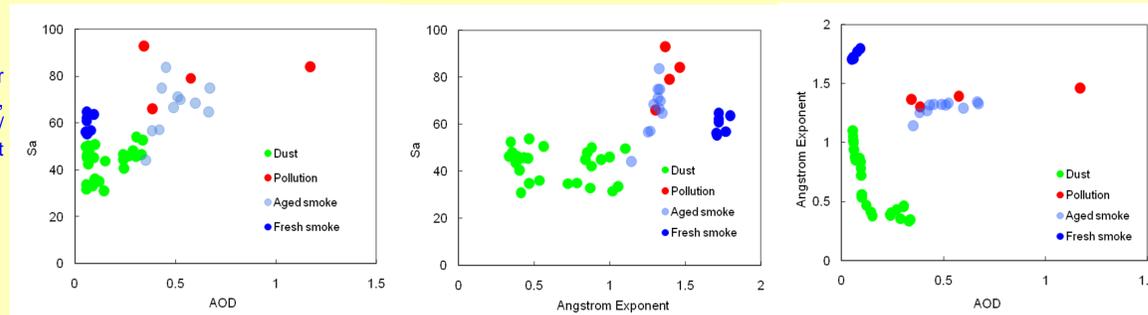
## Extinction-to-backscatter ratios ( $S_a$ ) for key aerosol types (linking the sunphotometry and the lidar profiles)

The lidar (extinction-to-backscatter) ratio is the key to linking sunphotometer AODs with integrated profiles of lidar backscatter coefficient (the next step in our analysis). They were computed for key aerosol types of dust, aged smoke, fresh smoke and pollution events at  $0.5 \mu\text{m}$  on the basis of selected retrievals of aerosol properties from AEROCAN/AERONET over Egbert. The Lidar ratio has been calculated from the following equation (Muller et al., 2003; Dubovik et al., 2006)

$$S_a(\lambda) = \frac{4\pi}{\omega_{0,a} p_a(\cos \pi)}$$

• Low mean values of the lidar ratio,  $S_a$ , at  $0.5 \mu\text{m}$  for desert dust ( $\langle S_a \rangle \approx 43 \text{ sr}$ ) aerosols are clearly distinguishable from biomass burning (Aged smoke:  $\langle S_a \rangle \approx 67 \text{ sr}$  and Fresh smoke  $\langle S_a \rangle \approx 60 \text{ sr}$ ) aerosols and pollution ( $\langle S_a \rangle \approx 81 \text{ sr}$ ).

where  $p_a(\cos \pi)$  is the backscattering value of the aerosol phase function.  $\omega_{0,a}$  is the single-scattering albedo.  $p_a$  at  $500\text{nm}$  was calculated through the linear interpolation of Dubovik inversion values at  $440$  and  $676\text{nm}$ .



• Our next step will be to investigate the utility of further dividing the classification scheme into coarse and fine mode domains.

Table 1: Summary of Aerosol Optical Parameters for key aerosol types (Retrieved From AEROCAN/AERONET)

| Type        | $S_a$    | $\alpha$  | AOD       | Source Region   |
|-------------|----------|-----------|-----------|---|
| Dust        | 42.89±7  | 0.65±0.26 | 0.16±0.1  | Asia  |
| Pollution   | 80.58±11 | 1.38±0.07 | 0.62±0.38 | Regional smog   |
| Aged Smoke  | 66.59±11 | 1.30±0.06 | 0.5±0.1   | Forest fires in Siberia                               |
| Fresh Smoke | 59.95±4  | 1.73±0.03 | 0.07±0.02 | Relatively fresh smoke aerosols from northern Ontario |

• The results indicate that  $S_a$  can be coarsely classified using the Angstrom exponent and the AOD.

## Conclusions

- A combination of ground-based measurements coupled with remote sensing imagery products and back trajectories were used to characterize a variety of aerosol events.
- This yielded contributions to a climatology of  $S_a$  values (the essential link between sunphotometry and integrated lidar backscatter coefficient profiles)
- Long term goal: a measure of the variability in  $S_a$  for key aerosol types would be useful for inverting global lidar observations (notably CALIOP).